

Taphonomy, sedimentology and chronology of a fossiliferous outcrop from the continental Pleistocene of Uruguay

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RESUMEN

La fauna local La Paz (34°44' S, 56°17' W, elevación 31 m) constituye una asociación fosilífera con abundantes restos de mamíferos típicos del Pleistoceno del sur de Sudamérica. Esta asociación se encuentra conformando un banco osario de características únicas en Uruguay, el cual contiene una gran concentración de restos óseos en su mayoría completamente desarticulados y fragmentados, aunque también otros, que están muy bien conservados. Los principales objetivos de este trabajo son contribuir al entendimiento de los procesos sedimentarios y tafonómicos actuantes y aportar nuevos datos a la cronología del depósito. Los fragmentos óseos estudiados se presentan sin conexión anatómica, en su mayoría sumamente fracturados, y acumulados conformando un banco osario multidominante. El número de elementos identificados (NISP) es de 130. El 50% de las piezas estudiadas no presentó signos de meteorización; el 70 % mostró algún tipo de fractura, en la mayoría de los casos (40%) paralelas al eje principal; un escaso número de piezas mostró señales de abrasión (5%) y la evidencia directa indica que no existió selección (ni por forma ni por tamaño) por parte del agente de transporte. Los aspectos facirológicos, como la ausencia de estructuras sedimentarias -depósito predominantemente masivo-, la muy pobre grano-selección (clasificación), el elevado grado de angulosidad de los clastos mayores y las asociaciones de facies presentes permiten proponer que el fenómeno responsable del depósito final de los restos fue del tipo flujo de barro (mudflow) no canalizado. Las características tafonómicas también son compatibles con un proceso de este tipo. Los sedimentos portadores del banco osario son atribuibles a la Formación Libertad, asignada al Pleistoceno inferior-medio en la literatura precedente. Adicionalmente, se discute la validez de las dos únicas dataciones radiocarbónicas exitosas en el contexto de numerosos análisis fallidos. Los mamíferos de la fauna local La Paz y la presencia de ciertos minerales arcillosos permiten inferir, en el contexto de este depósito, un ambiente árido a semiárido.

Palabras clave: Cuaternario, banco osario, tafonomía, Formación Libertad, Uruguay.

ABSTRACT

The La Paz local fauna (34°44' S, 56°17' W, elevation 31 m) is a fossiliferous assemblage with abundant remains of typical southern South American Pleistocene mammals. The assemblage consists of a bonebed, containing a high concentration of skeletal remains, most of them completely disarticulated and fragmented, although very well preserved, and is unique in Uruguay as most remains have been found as single isolated finds until now. The goals of this study are to determine the sedimentary and taphonomic processes that produced this regionally unique concentration of bones and determine the age of the deposit. The bone fragments studied are not in their original anatomical position and consist

of mostly, very fractured specimens accumulated in a multidominant bonebed. The number of identified specimens (NISP) is 130. Within the recovered sample, 50% of the studied specimens showed no signs of weathering; 70% showed some kind of fracture or breakage, many of the specimens (40%) parallel to the major axis; and a small number of specimens showed signs of abrasion (5%). Based on the sample there does not appear to have been selection (either by form or by size) by transport. The dominantly massive deposit lacks sedimentary structures, has poor sorting and a high degree of angularity of the largest pebbles suggesting that the final deposition of the remains resulted from a mudflow and were not channeled deposited. The taphonomic features observed on the bones are also compatible with this depositional process. The sediments in which the bonebed is located are attributed to the Libertad Formation, currently assigned to the lower-middle Pleistocene by other workers. We discuss the validity of the two available radiocarbon ages in the context of many failed essays. The mammals of the La Paz local fauna and the presence of particular clay mineral allow us to infer, in the context of this deposit, an arid to semi-arid environment.

Key words: Quaternary, bonebed, taphonomy, Libertad Formation, Uruguay.

INTRODUCTION

The La Paz local fauna (34°44' S, 56°17' W, elevation 31 m, Figure 1) is a fossiliferous assemblage with abundant remains of mammals typical of the Pleistocene in southern South America. This fauna was obtained from a preserved bonebed in a diamictite facies of the Libertad Formation (Pleistocene *sensu lato*) stratigraphically above a weathered granular layer of the La Paz Granite (Cambrian) located in a quarry.

Preliminary studies of the faunal composition of the association, its chronology, paleoenvironments and processes related to the genesis of the deposit have been previously discussed (Perea *et al.*, 2001; Marchesano *et al.*, 2002), whereas a detailed study of the site was initiated in 2005. This site is unique in Uruguay, as it contains a high concentration of skeletal remains, most of them completely disarticulated and fragmented, although other part is very well preserved. The fauna includes characteristic taxa of the Pleistocene of southern South America: *Glyptodon* sp., *Doedicurus* sp., *Panochthus* sp., *Neosclerocalyptus* sp., *Lestodon* sp., *Macrauchenia* cf. *M. patachonica*, *Toxodon* sp., *Antifer* sp., *Stegomastodon waringi*, among others. However, our analysis of the taphonomy of the deposit suggests that the constituent elements of the assemblage do not represent a single community.

Based on the taphonomic and sedimentological evidence (Perea *et al.*, 2001; Marchesano *et al.*, 2002; Corona *et al.*, 2007; Corona *et al.*, 2008a) the bonebed appears to have formed as a result of transport and deposition of the faunal remains during a high-energy event, like a mudflow. In addition, the presence of sedimentary facies that indicates remobilization and aeolic reworking in part of the sedimentary succession, show the occurrence of successive cycles of transport of the bones and sedimentation in a continental environment.

Based on geological mapping of the area, it was initially proposed that the lithostratigraphic unit that includes the bonebed was the Dolores Formation (Late Pleistocene) (Perea *et al.*, 2001; Marchesano *et al.*, 2002). However, Corona *et al.* (2008a) considered that the bonebed, most

likely, is located in the Libertad Formation taking into account that the studied section occupies upper topographic positions, the slightly undulating morphology, and its location on the 1:100000 scale geological map of the area in which the bonebed occurs (Coronel *et al.*, 1988). In addition, Gutiérrez *et al.* (2005) obtained a radiocarbon age of 17620 ± 100 yrs BP for this local fauna, based on dental enamel of *Stegomastodon waringi*, while here we present new chronological data.

The main objectives of this study are to contribute to the understanding of taphonomical and sedimentary processes that produced this unique concentration

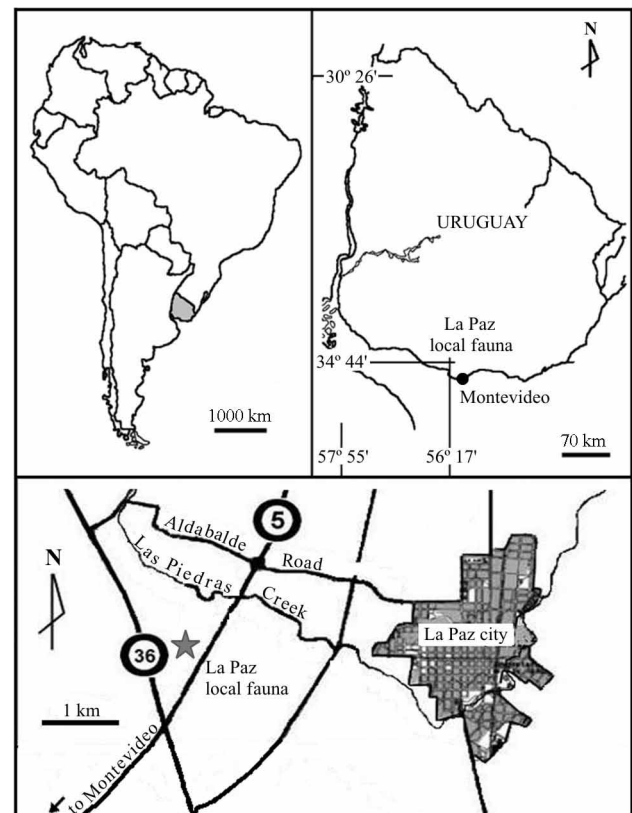


Figure 1. Location of the studied fossil outcrop (La Paz local fauna).

of bones never before documented in Uruguay, as well as to provide new data regarding the age of the deposit. Specifically, we made a comprehensive analysis of the taphocenosis through its main parameters. We investigated the sedimentology of the section in order to identify various lithofacies and analyzed the mineralogy of the clay fraction. Additionally, we analyze and discuss some chronological aspects.

METHODS

Taphonomy

Following the procedures described in Beherensmeyer (1978), Lyman (1994) and Cladera *et al.* (2004), we examined the following taphonomical parameters of the bones: orientation, degree of disarticulation, weathering stages, amount of abrasion, fracturing and degree of selection of the remains. The first two parameters were studied *in situ*, while the others were determined in the laboratory following the removal of the bones from the bonebed.

Orientation

It means the position of the long axis of bones to determine if there is any preferential direction.

Disarticulation

It refers to the degree to which the original anatomical connections of skeletal parts were preserved. We distinguished: elements retaining their original anatomical connection and elements without anatomical connections.

Weathering

It involves the loss of material from bone surface as result of physical and chemical agents. Beherensmeyer (1978) recognized six states of the weathering process, but for our study we simplified this to three categories: 1 (intact bone, corresponds to the state 0 of Beherensmeyer); 2 (loss of superficial material, corresponding to Beherensmeyer's 1 and 2 stages) and 3 (loss of deep bone material, Beherensmeyer's stages from 3 to 5).

Fracturing

It refers to both the number and disposition of fractures present in the remains. It was established using the following categories for both, the number and kind of fracture: 0 (no fracture), 1 (one or more fractures across the maximum dimension), 2 (one or more oblique fractures to the maximum dimension); >2 (one or more fractures longitudinal to the maximum dimension).

Abrasion

It measures the softness of edges and loss of roughness by transport. The categories established were: 1 (intact bone), 2 (rounded bone), 3 (polished bone).

Selection

First was established the shape of the elements following the Zingg's criterion for gravel (Suguio, 2003), which can be applied considering fossils as bioclasts. According to the relations among the three axes, Zingg established four basic forms: discoid, spherical, cylindrical and ellipsoidal. The three axes of each specimen (a, b, c) perpendicular to each other, being $a > b > c$ were measured and the data plotted similar to Zingg's diagram.

Simultaneously, we identified taxonomically and anatomically each specimen and established the age of the individuals (young, young adult, adult and senile) based on the stages of tooth eruption or wear of the dentition or for the post-cranial skeleton through size or degree of ossification-of other elements. Some of the most important specimens in terms of better preservation are described in "Systematic Paleontology".

The studied fossil materials (Table 1) are housed in the Colección Paleontológica de la Facultad de Ciencias, Montevideo, Uruguay (FC-DPV) and in the Colección Particular Luis Castiglioni (LRC).

Sedimentology

For the studied outcrop we built a detailed stratigraphic profile in which textures, sedimentary structures, geometry and fossiliferous content were indicated. Representative samples of each lithofacies were collected in order to carry out mineralogical studies. Samples were subjected to physical and chemical breakdown for separation by screening the clay fraction. Clay components were treated with conventional methods for dispersion (sodium hexametaphosphate) in order to analyze by X-ray diffraction. These tests were carried out at the Instituto de Física, Facultad de Ciencias, with a diffractometer PW3710 (Anode: copper, wavelength = 1.54056 Å). The sedimentological study was complemented with textural and compositional observation of the coarser fraction (sand and silt) through a binocular magnifier and petrographic microscope.

Chronology

Some of the better preserved specimens were selected to obtain radiocarbon ages. These analyses were performed in the Accelerator Mass Spectrometry Laboratory, University of Arizona and in LATYR, Universidad Nacional de La Plata.

SYSTEMATIC PALEONTOLOGY

Included here are brief descriptions of some of the most relevant materials in terms of better preservation (Figure 2).

Mammalia Linnaeus, 1758
 Xenarthra Cope, 1889
 Folivora Delsuc *et al.*, 2001
 Mylodontidae Gill, 1872
Lestodon Gervais, 1855
***Lestodon* sp.**

Materials. LRC 621, right femur. LRC 838, almost complete skull with damaged anterior dorsal region.

Measurements. LRC 621. Length: 74 cm; width: 24 cm.

Comments. Both specimens are very large representatives of the genus.

Notoungulata Roth, 1903
 Toxodontidae Gervais, 1847
Toxodon Owen, 1837
***Toxodon* sp.**

Material. LRC 469. Almost complete skull lacking incisors.

Measurements. Length: 53.5 cm; width: 35.8 cm.

Comments. It seems to be a young specimen because of the small size and the porous texture of some bones.

Carnivora Bowdich, 1821
 Ursidae Gray, 1825
 Tremarctinae Merriam y Stock, 1925
Arctotherium Bravard, 1857
***Arctotherium bonariense* Gervais, 1852**

Material. LRC 740, right mandibular fragment with p4, damaged m1 and m2.

Measurements. Teeth row length: 9.8 cm; mandible depth between m1 and m2: 7.3 cm; p4: 1.9/2.2 cm; m1: 2.6/4.5 cm; m2: 2.3/3.6 cm.

Comments. The teeth morphology allows us to recognize the species, being this specimen a very large representative of it.

RESULTS

Taphonomy

The studied bone fragments that constitutes the single bonebed are disarticulated from their original anatomical positions, most are highly fractured, and accumulated in a chaotic way with no preferred orientation and densely packed. Overall, this bonebed, which shows lobular geometry, has a lateral continuity of about 50 m and a maximum thickness of 1.0 m. The highest density of specimens is found in the central area, decreasing significantly towards the marginal areas, from where the best preserved fossils (skulls, jaws and the most complete long bones) were exhumed (Figure 3).

According to Rogers *et al.* (2007) this is a macrofossil bonebed (most bioclasts are more than 5 cm in maximum

dimension). Also, it is a high diversity bonebed (10 or more taxa), and multidominant, since more than two taxa contribute to the total number of elements identified (NISP). Assuming that the probability of association among the elements is zero due to the high degree of disarticulation observed and the possible physical and/or temporary mixture, it is possible to obtain the number of individuals from the NISP (*sensu* Rogers *et al.*, 2007). We obtained a minimum of 130 individuals, mostly representing herbivorous and omnivorous forms. The carnivores are represented by the short-faced bear *Arctotherium bonariense* (Corona *et al.*, 2008b) and a phorusracid bird (Alvarenga *et al.*, 2010). It should be noted that, due to the very fragmentary nature of the specimens, many could not be anatomical and/or taxonomically identified, therefore they were not included in the quantitative taphonomic analysis.

We identified two taphonomic modes (*sensu* Beherensmeyer and Hook, 1992) within the bonebed: the first and most frequent consists of highly fragmented and fragile bones with a characteristic yellowish white color, usually with dark spots on the surface resulting from mineral precipitation during diagenesis. The second is formed by very consolidated, black, small bones with rounded vertices and edges, indicating longer transport time and possible selection by size. Many of the bone specimens have gravel-size clasts adhering to its surface, with calcium carbonate either in the form of crusts (concretions) or as cement. The crusts and carbonate nodules may be formed after deposition during burial, mainly during early diagenetic time. It looks like randomly dispersed over the section, but they are mainly concentrated or nucleated around the bonebed. This indicates the early time diagenetic dissolution and precipitation of calcite around fossils. In addition, many of the fossils provide evidence of relatively long periods of exposure to erosive agents, while others have intact surfaces.

We describe below the main results of the taphonomic analysis (Figure 4):

Weathering

Fifty percent of the specimens did not show weathering effects (stage 1), while 40% showed loss of superficial bone (stage 2), but not affected the entire bone surface evenly. Only 10% of the materials were more affected, with loss of part of spongy bone (stage 3).

Fracturing

Seventy percent of the specimens had some type of fracture. In 40% of the cases the fractures are parallel to the main axis. Transverse and oblique fractures occurred in less proportion, 20% and 10% respectively.

Abrasion

The majority of fossils (95%) showed no evidence of abrasion. A very small percentage (5%) showed a marked erosion of reliefs and edges, attributed to category 2. There were no materials attributable to category 3 (polished bone).

Table 1. List of materials used in the taphonomic analysis.

Collection ID	Material	Taxa
LRC 233	phalanx	Xenarthra
LRC 374	navicular?	Xenarthra
FC-DPV 1669	phalanx	Pilosa
FC-DPV 1671	metapodial	Pilosa
FC-DPV 1268	metapodial	Pilosa
FC-DPV s/n	vertebrae	Mylodontidae
FC-DPV s/n	vertebrae	Mylodontidae
LRC 395	caudal vertebrae	Mylodontidae
LRC 395	caudal vertebrae	Mylodontidae
LRC 562	phalanx	Mylodontidae
LRC 436	phalanx	Mylodontidae
LRC 568	fragment of skull	Mylodontidae (senile)
LRC 624	dorsal vertebrae	?Mylodontidae
LRC 624	dorsal vertebrae	?Mylodontidae (young)
FC-DPV 1812	fragment of mandible with m2-m4	<i>Lestodon</i>
FC-DPV 2127	hand bone	<i>Lestodon</i>
FC-DPV 2149	vertebra	<i>Lestodon</i>
LRC 221	fragment of mandible	<i>Lestodon</i> (adult)
FC-DPV 1375	astragalus	<i>Lestodon</i>
FC-DPV 1656	almost complete tibia	<i>Lestodon</i>
LRC 236	almost complete tibia	<i>Lestodon</i>
LRC 362	cuboid?	<i>Lestodon</i>
LRC 458	phalanx	<i>Lestodon</i>
LRC 460	tarsal bone	<i>Lestodon</i>
LRC 552	fragment of mandible	<i>Lestodon</i> (young)
LRC 561	fragment of mandible	<i>Lestodon</i>
LRC 572	almost complete tibia	<i>Lestodon</i> (young)
LRC 574	almost complete tibia	<i>Lestodon</i>
LRC 588	fragment of humerus	<i>Lestodon</i>
LRC 615	calcaneus	<i>Lestodon</i>
LRC 618	astragalus	<i>Lestodon</i>
LRC 619	almost complete mandible ramus	<i>Lestodon</i>
LRC 621	almost complete femur	<i>Lestodon</i>
LRC 643	sesamoid?	<i>Lestodon</i>
LRC 623	molariform	Megatheriinae
FC-DPV 1666	carpal or tarsal bone	Glyptodontidae
FC-DPV 1668	carpal or tarsal bone	?Glyptodontidae
FC-DPV 1670	phalanx	?Glyptodontidae
FC-DPV 1672	phalanx	?Glyptodontidae
LRC 617	astragalus	Glyptodontidae
FC-DPV 1269	osteoderm	<i>Glyptodon</i>
FC-DPV 2129	osteoderm	<i>Glyptodon</i>
FC-DPV 2130	osteoderm	<i>Glyptodon</i>
FC-DPV 2131	osteoderm	<i>Glyptodon</i>
FC-DPV 2132	osteoderm	<i>Glyptodon</i>
FC-DPV 2133	osteoderm	<i>Glyptodon</i>
FC-DPV 2134	osteoderm	<i>Glyptodon</i>
FC-DPV 2135	osteoderm	<i>Glyptodon</i>
FC-DPV 2136	osteoderm	<i>Glyptodon</i>
FC-DPV 2137	osteoderm	<i>Glyptodon</i>
FC-DPV 2138	osteoderm	<i>Glyptodon</i>
FC-DPV 2139	osteoderm	<i>Glyptodon</i>
FC-DPV 2140	osteoderm	<i>Glyptodon</i>

continues

Table 1 (continued). List of materials used in the taphonomic analysis.

Collection ID	Material	Taxa
FC-DPV 2141	osteoderm	<i>Glyptodon</i>
LRC 224	osteoderm	<i>Glyptodon</i>
LRC 372	osteoderm	<i>Glyptodon</i>
LRC 372	osteoderm	<i>Glyptodon</i>
LRC 372	osteoderm	<i>Glyptodon</i>
LRC 372	osteoderm	<i>Glyptodon</i>
LRC 408	osteoderm	<i>Glyptodon</i>
LRC 408	osteoderm	<i>Glyptodon</i>
LRC 408	osteoderm	? <i>Glyptodon</i>
FC-DPV 1373	dorsal osteoderm	<i>Panochthus</i>
FC-DPV 2142	osteoderm	<i>Panochthus</i>
FC-DPV 2143	osteoderm	<i>Panochthus</i>
FC-DPV 2144	osteoderm	<i>Panochthus</i>
FC-DPV 2145	osteoderm	<i>Panochthus</i>
LRC 408	osteoderm	<i>Panochthus</i>
LRC 408	osteoderm	<i>Panochthus</i>
LRC 682	carapace fragment	<i>Doedicurus</i>
LRC 408	osteoderm	<i>Neosclerocalyptus</i>
FC-DPV 1374	dorsal osteoderm	<i>Neosclerocalyptus</i>
LRC 672	almost complete humerus	Pampatheriidae
FC-DPV 1371	osteoderm	<i>Propraopus</i>
FC-DPV 1263	premaxilla with incisors	Toxodontidae
FC-DPV 1270	molar fragment	Toxodontidae
FC-DPV 1271	fragment of maxilla	Toxodontidae (young)
FC-DPV 1275	premaxilla with alveolus	Toxodontidae
FC-DPV 1665	fragment of mandible with m2 y m3	Toxodontidae (young)
FC-DPV 2148	forelimb bone	Toxodontidae
LRC 391	forelimb bone	Toxodontidae
LRC 575	almost complete tibia	Toxodontidae
LRC 225	fragment of mandible	<i>Toxodon</i>
LRC 220	almost complete mandible	<i>Toxodon</i>
LRC 369	fragment of premaxilla with teeth	<i>Toxodon</i>
LRC 469	almost complete skull	<i>Toxodon</i>
LRC 625	fragment of mandible	<i>Toxodon</i>
LRC 718	fragment of mandible with teeth	<i>Toxodon</i>
LRC 400	molar	Artiodactyla (young)
LRC 400	molar	Artiodactyla (young)
FC-DPV 1273	fragment of mandible with teeth	Cervidae (young adult)
FC-DPV 1276	fragment of mandible with teeth	Cervidae (young adult)
LRC 376	astragalus	Cervidae
LRC 376	astragalus	Cervidae
LRC 564	astragalus	Cervidae
LRC 379	hand bone	Cervidae?
LRC 407	fragment of radius	Cervidae
LRC 404	antler fragment	Cervidae
LRC 404	antler fragment	Cervidae
LRC 404	antler fragment	Cervidae
LRC 404	antler fragment	Cervidae
LRC 404	antler fragment	Cervidae
LRC 404	antler fragment	Cervidae
LRC 539	metapodial	Cervidae
FC-DPV 1274	fragment of mandible	Camelidae (young adult)

continues

Table 1 (continued). List of materials used in the taphonomic analysis.

Collection ID	Material	Taxa
FC-DPV 1376	almost complete metatarsus	Camelidae
FC-DPV 1856	proximal phalanx	Camelidae
FC-DPV 1857	metapodial distal fragment	Camelidae
FC-DPV 2128	distal fragment of metapodial	Camelidae
LRC 462	fragment of maxilla with teeth	?Camelidae (young adult)
LRC 544	fragment of mandible	?Camelidae (young adult)
LRC 622	fragment of maxilla with teeth	?Camelidae (senile)
LRC 674	proximal fragment of metapodial	Camelidae
LRC 712	Proximal end of metapodial	?Camelidae
FC-DPV 1370	almost complete mandible	<i>Stegomastodon waringi</i>
FC-DPV 1803	fragment of maxilla with M3	<i>Stegomastodon waringi</i>
LRC 219	fragment of tusk	Gomphotheridae
LRC 219	fragment of tusk	Gomphotheridae
LRC 456	navicular?	Gomphotheridae
FC-DPV 1654	lower molar	<i>Hippidion</i>
LRC 373	metapodial	<i>Equus</i>
LRC 457	astragalus	<i>Equus</i>
FC-DPV 1272	fragment of mandible with teeth	Macrauchenidae (young adult)
FC-DPV 1655	fragment of mandible with m2 and m3	Macrauchenidae (young adult)
LRC s/n	claw	Aves
LRC 677	fragment of tarsometatarsus	Phorusracidae

LRC: Colección particular Luis Castiglioni, Montevideo, Uruguay.
FC-DPV: Colección Paleontológica, Facultad de Ciencias, Montevideo, Uruguay.

Selection

There was no sorting during transport, as included in the mud debris deposit are both small (about 3 cm, *e. g.* Phorusracidae tarsometatarsus) and large (80 cm, *e. g.* *Lestodon* femur). As shown in Figure 4d, all the studied elements were distributed randomly in the Zingg diagram, which implies that there was no selection by shape.

Sedimentology

The complete stratigraphic profile (Figure 3) allowed the recognition of distinct depositional events. The studied section shows three lithofacies.

Lithofacies 1

Deposits of diamictite and mud with relatively well-defined, thin horizontal lamination, rest unconformably on the La Paz Granite (Cambrian). These sediments are dark grey-brown (10YR 4/2), 2 m thick, have a tabular geometry and consist of silty claystone with black spots, with very scattered angular clasts from the granite below, that varies in size from gravel to fine-coarse sand made of quartz, feldspar, and biotite. Associated with this level are

rizholiths, concretions and thick crusts of calcium carbonate. Toward the top are interspersed lenses of thin, green clay.

Lithofacies 2

In abrupt contact above the previous lithofacies is a white (10 YR 8/1), matrix-supported diamictite, with relatively tabular geometry, and variable thickness between 2.0 and 2.50 m. Very angular megaclasts made of quartz, are abundant; they vary in size between gravel and pebble. The matrix is predominantly clayey containing a few sheets of biotite. At the bottom of this lithofacies the bonebed is found and consists of abundant remains of Pleistocene mammals, which show a relatively large lateral distribution (tens of meters). Above the bonebed appears a level of hard crust composed of calcium carbonate.

Lithofacies 3

The contact with the underlying unit is abrupt. It consists of a brown (7.5YR 5/4), approximately 1.50 m thick, sludge (predominantly silt clayey fraction) with tabular geometry, massive to slightly laminated structure and some oxide spots. This sludge includes some well-sorted, fine to medium sand composed of rounded to subrounded quartz.

The results of X-ray diffraction of samples of the three lithofacies show a predominance of chlorite and smectite over sepiolite. In particular, the lithofacies 2 (where is the bonebed) contains chlorite and sepiolite.

Chronology

The number of skeletal elements of this association is high, but the fragmentary nature of most of them do not allow a determination to a specific level (Appendix). However, the bonebed does not consist of taxa with stratigraphic or chronologic range restricted to a single Quaternary chronostratigraphic unit (*sensu* Cione y Tonni, 2005). *Stegomastodon waringi* and *Arctotherium bonariense* are restricted to the Bonaerian-Lujanian stages (late middle Pleistocene-early Holocene). It is noteworthy that from dental enamel of *Toxodon* a new age was obtained for this local fauna: 5917 ± 65 years BP. This one and that performed by Gutiérrez *et al.* (2005) were the only data obtained in the context of a lot of failed analysis and cannot be considered indicative of the true age of the bonebed (Table 2).

DISCUSSION

Taphonomy

The taphonomic features of the bonebed are indicative of a deposit generated by a high energy process, such as a mudflow. The degree of weathering observed in the fossils (only 50% of materials analyzed) indicates that exposure prior to burial was long enough to allow complete disarticu-

lation and dispersal of skeletal components, but not long enough to produce deep weathering signs.

However, as shown above, note that most of the fossils presented fractures parallel to the main axis of the specimens. According to Haynes (1980) and Alcalá (1994) fractures of this type are due to processes such as contraction and expansion of the bones during biostratinomic stage, prior to final burial. Therefore, we could infer only from this result that the subaerial exposure period was quite long, which is not consistent with the previous observation on the lack of deep weathering of the bone surface. According to Merino-Cristóbal (2000), fractures that occur prior to burial show cracks filled of sediment or some kind of cement, which was observed in many of the fossils from the La Paz local fauna. Transverse and oblique fractures occurred in a smaller number of specimens. The first fracturing event,

presumably, occurred during diagenesis (Fernández-López, 2000), while there is no agreement among researchers on the factors responsible for the second event: action of carnivores, fluvial transport or trampling have been proposed (Cladera *et al.*, 2004).

The observation and analysis of degree of weathering and type of fractures are mutually inconsistent results: the first parameter indicates a short time of subaerial exposure; the second suggests otherwise. However, the effects of weathering of the bone can be mitigated if the remains are under conditions of high humidity and protected by abundant vegetation (Behrensmeier, 1978) which, in the current context cannot be excluded. Moreover, the presence of various stages of weathering may also indicate the occurrence of more than one taphonomic history, either by accumulation over long periods or concentration under different condi-

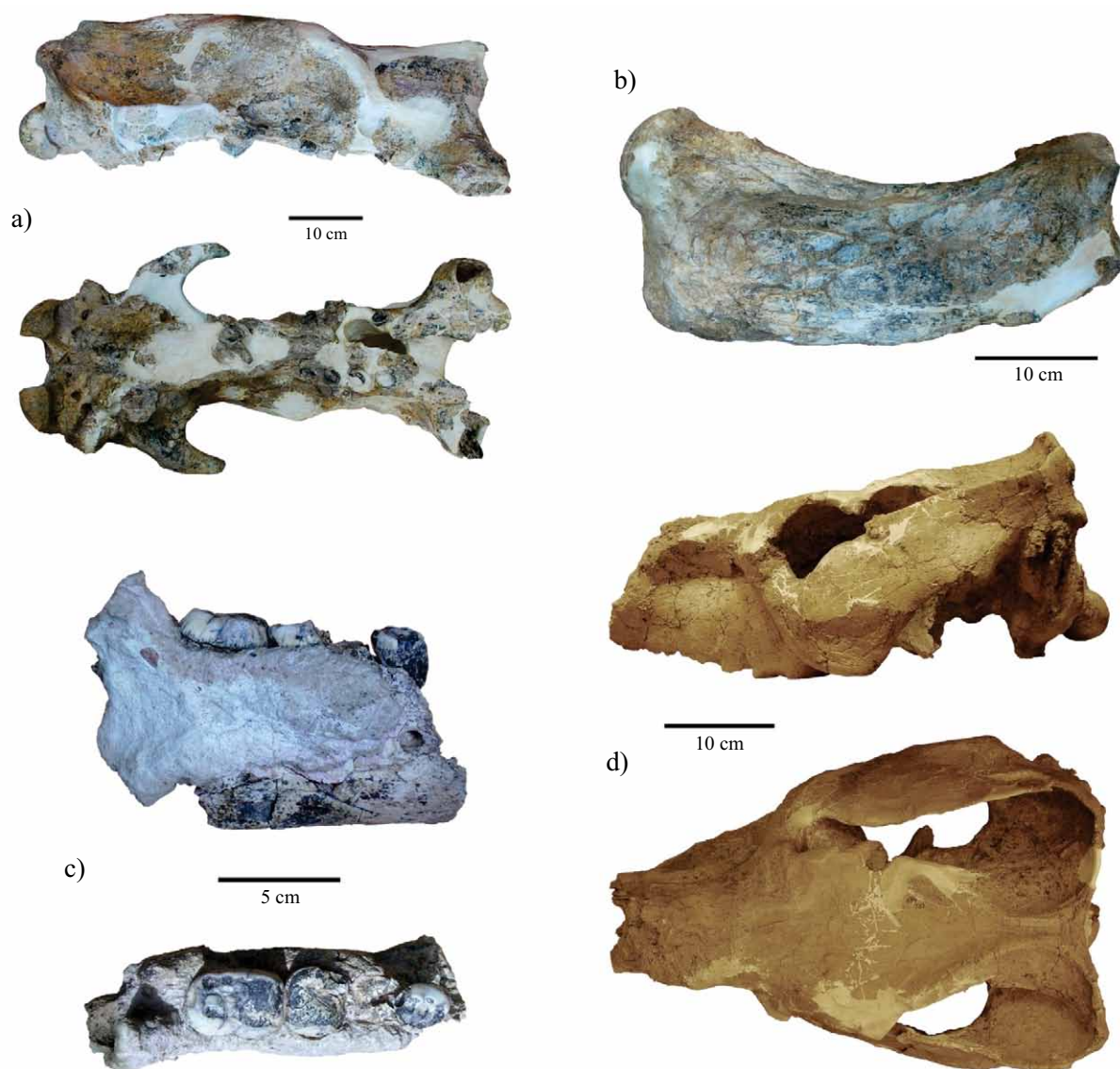


Figure 2. Some of the most relevant materials described above: a) skull of *Lestodon* sp. in lateral and palatal views; b) femur of *Lestodon* sp.; c) ramus of mandible of *Arctotherium bonariense* in lateral and occlusal views; d) skull of *Toxodon* sp. in lateral and dorsal views.

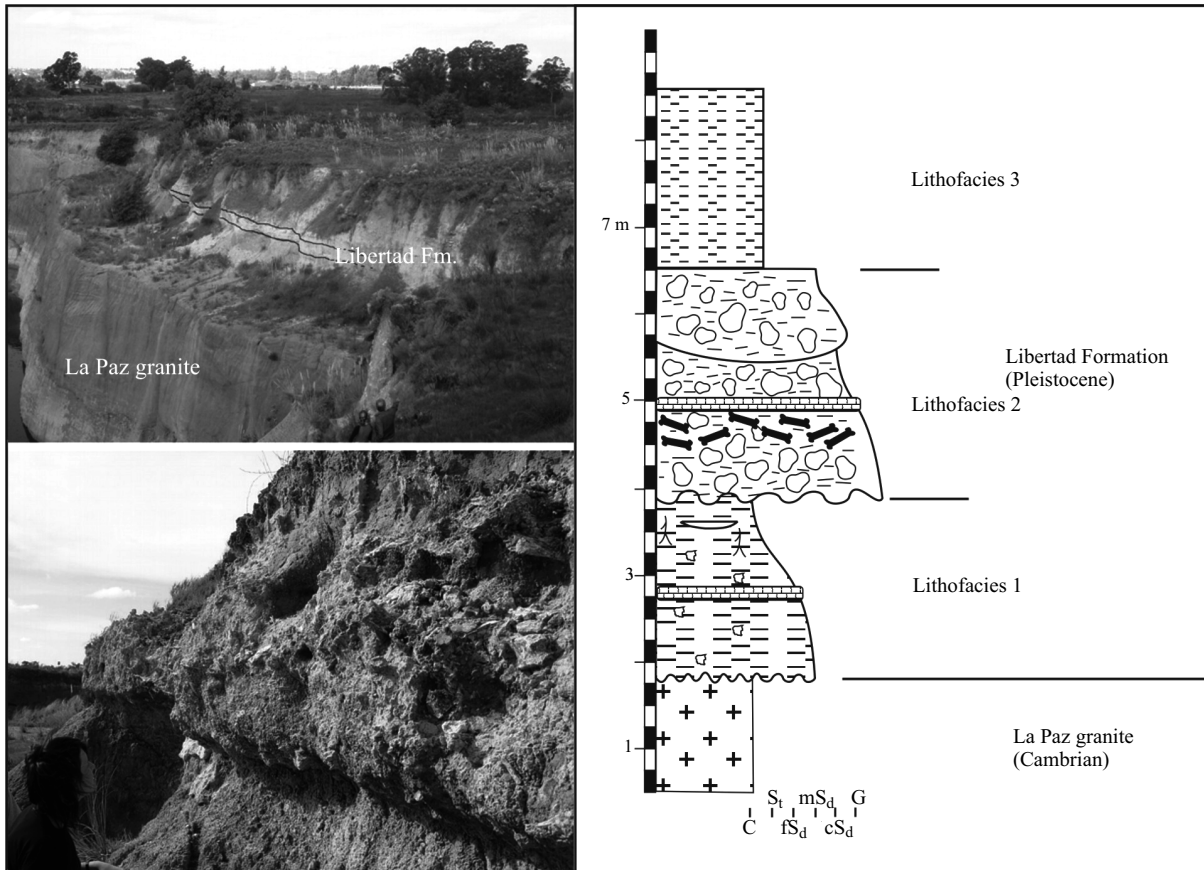


Figure 3. Panoramic view of bonebed and details of the central area where the greater accumulation of elements occurs. To the right is a detailed stratigraphic profile.

tions after the initial weathering (Behereismeyer, 1978; Rogers *et al.*, 2007).

The small number of specimens that exhibit abrasion (5%) indicates that the impact on these materials, or other debris making up the load, was low. This observation, together with direct evidence that there was no selection (either by shape or by size) during transport, is compatible with the hypothesis of the occurrence of a high energy event responsible for the deposition. Furthermore, the characteristics of the bonebed in terms of massive and chaotic arrangement of its components and the presence of fossils from different backgrounds are consistent with a high energy gravity flow. The mixture of fossils with different taphonomic histories indicates that this assemblage is parautochthonous (remains reworked in the depositional site) or allochthonous (formed by remains transported from other areas) or a combination of both.

Sedimentology

The characteristics of the diamictite like an internal massive bedding or absence of sedimentary structures, poor sorting, erosive contact at the base, high angularity

of the largest clasts, the bioclast supported texture, the tabular extent geometry and, most important, the facies associations and the continental environment involved, allow to suggest that the final deposition occurred by non-channelized, subaerial mudflow processes. The lithology and facies arrangement are consistent with that described by Hauser (1985) and Lopes Da Silva *et al.* (1997). This type of process permits to gravel clasts (or in this case bioclasts or bones) to move in a cohesive, very thick matrix of clay, silt and water. These flows result in the generation of massive deposits, without structures and with poor sorted clasts, scattered throughout and supported by a pelitic matrix (Major, 2003). At present this type of deposits are typically generated by mudflows in alluvial fans and alluvial environments, especially when the discharge is significant, due to abundant rainfall (Eberth *et al.*, 2006).

Lithostratigraphy

With regard to the designation of the bonebed to a distinct lithostratigraphic unit, it was recently proposed that the sediment containing the La Paz local fauna is part of the Libertad Formation (Corona *et al.*, 2008a). However, this

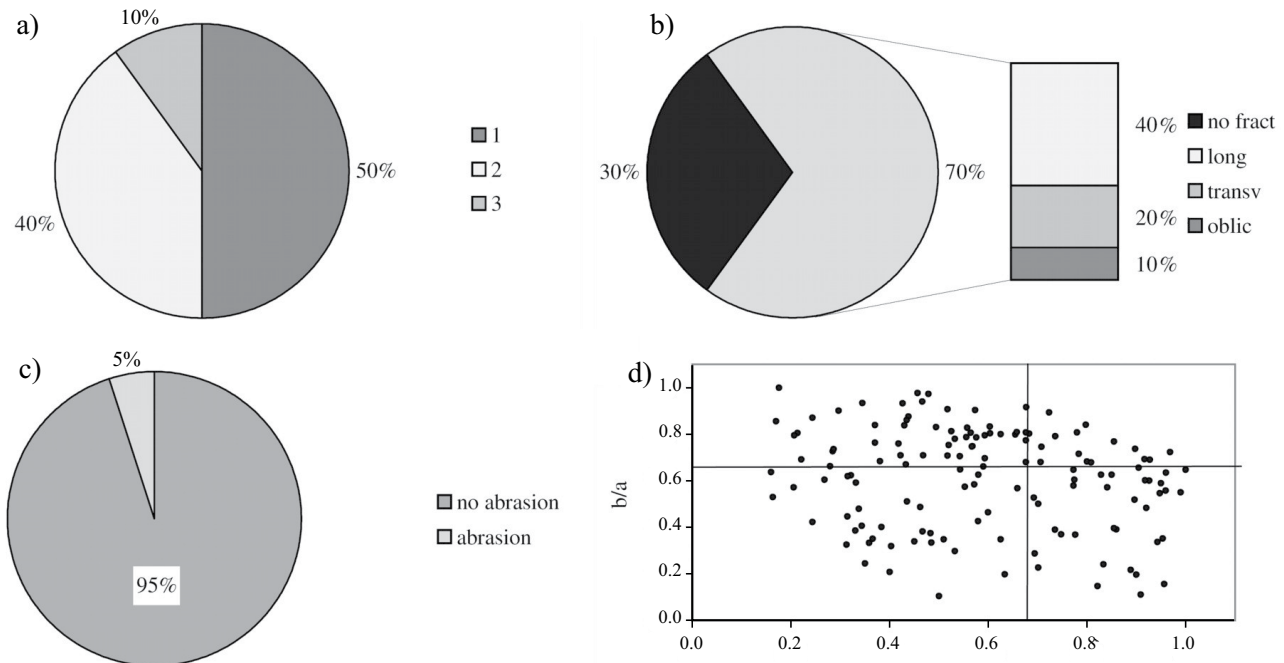


Figure 4. Results of quantitative analysis of taphonomic parameters. a) weathering; b) fragmentation; c) abrasion; d) selection.

assignment has varied depending on the author. Marchesano *et al.* (2002) proposed that the sedimentary strata which includes this local fauna belongs to Dolores Formation. Subsequently, we used a more informal name, designating these deposits as “diamictites sediment of Libertad-Dolores(?) Formation” (Corona *et al.*, 2007).

More generally, different authors have had difficulty assigning a formal name to the lithostratigraphic unit (Dolores Formation or Libertad Formation) for the Quaternary fossiliferous deposits in the southern part of Uruguay. The Dolores Formation (*sensu* Goso, 1972) is characterized by brown with gray-green local shades, silty clay, silt, clay, sand and gravel, associated with plains (Martínez and Ubilla, 2004). The Libertad Formation (*sensu* Goso, 1965) in contrast, consists of mudstones, sandy loess and reddish brown with calcium carbonate multiforms and undulate geomorphological features. This general lithologic similarity between both units is the cause of the problems to differentiate these two formations. In order to apply a consistent, formal designation, it essentially requires to apply geomorphologic criteria to distinguish them.

X-ray diffraction analysis of the clay fraction of the Quaternary continental deposits, confirmed the presence of smectite and illite in sediments of the Dolores Formation. On the other hand the Libertad Formation also contains these clays and, in addition, kaolinite (Goso and Ledesma, 2002; Goso *et al.*, 2002). The sediments containing the La Paz local fauna consist of smectite, chlorite, and sepiolite (Corona *et al.*, 2008a), which demonstrates that this methodology is not conclusive in the absence of major mineralogical differ-

ences between these two lithostratigraphic units. Therefore, although this technique contributes to the lithological characterization, for the moment does not distinguish between the two formations.

Chronology

Traditionally it has been assumed that the Libertad Formation was lower middle Pleistocene in age, although most of the previous works reaching this conclusion have not been stratigraphically controlled (Martínez y Ubilla, 2004). Although in the La Paz local fauna there are no taxa biostratigraphically restricted to a single Quaternary chronostratigraphic unit (*sensu* Cione y Tonni, 2005), the presence of *Stegomastodon waringi* and *Arctotherium bonariense* suggest a correlation with Bonaerian-Lujanian stages (late middle Pleistocene-early Holocene). Recently, substantial changes to the systematics of South American gomphotheres have been proposed by Mothé *et al.* (2011): *Stegomastodon* is a junior synonym of *Notiomastodon*, which includes all South American gomphotheres with a single species: *Notiomastodon platensis* (early Pleistocene-early Holocene). We agree with the conclusions of Alberdi and Prado (1995), Prado *et al.* (2005) and Alberdi *et al.* (2011), considering that *Stegomastodon waringi* a valid species which was distributed during middle to late Pleistocene in Ecuador, Brazil and Uruguay.

Unfortunately the sedimentary and genetic characteristics of this deposit does not allow the use of alternative

Table 2. List of data samples to obtain radiocarbon ages.

Lab Id	Material	Sample Id	F (d ¹³ C)	¹⁴ C age BP	± ¹⁴ C age
AA81956	tooth	<i>Toxodon</i>	0.4787	5917	65
AA86141	osteoderm	<i>Glyptodon</i>	x	x	x
AA86142	tooth	Cervidae	x	x	x
AA87517	tooth	<i>Toxodon</i>	x	x	x
AA87518	tooth	<i>Toxodon</i>	x	x	x
AA89360	tooth	<i>Toxodon</i>	x	x	x
AA89361	tooth	<i>Stegomastodon waringi</i>	x	x	x
1823	bone	Mammalia indet.	x	x	x

dating methods such as Optically Stimulated Luminescence or Thermoluminescence which had been widely used to date other Quaternary units. Therefore, one of the goals of our work was to obtain samples that could permit obtaining new radiocarbon ages. Collagen (organic matter present in a ratio of 18% in the bone) is rapidly degraded, often making it very difficult to obtain enough material for accurate dating (Gupta and Polach, 1985). Another major difficulty related to the dating of bones by the method of ¹⁴C is the potential presence of exogenous carbon contamination, either through organic molecules mobilized from the surrounding sediment or through metabolic products of microorganisms. Contamination is very difficult to detect (especially if diagenesis is severe and the quality of preservation of the materials is not good) and result in younger age than the true one (Hedges and Van Klinken, 1992; Arslanov and Svezhentsev, 1993). All samples from the La Paz local fauna, except one, did not preserve collagen (probably because it suffered extensive diagenesis), so they could not be dated. Following strict protocols, a new sample from the same specimen of *Stegomastodon waringi*, which was dated as 17620 ± 100 years BP (Gutiérrez *et al.*, 2005), was removed. The mandibular bone was perforated and a sample of 5 g of enamel was extracted from a not yet erupted tooth, superficially appeared unaltered and without fractures. However, the analysis determined that the sample lacked collagen, which contradicts the previous result.

An unexpected date of 5917 ± 65 years BP from dental enamel of *Toxodon*, a large difference with previously published age 17620 ± 100 years BP (Gutiérrez *et al.*, 2005) and the absence of collagen in the remaining samples (including a specimen previously dated), suggest that the validity of both absolute ages is highly questionable.

Paleoenvironment and paleoclimate

There is no agreement among authors regarding the type of climatic conditions that existed at the time of deposition of Libertad Formation. The ranges of suggested climatic conditions that have been proposed include wet periods,

torrential conditions and glacial to semi-arid episodes (Martínez and Ubilla, 2004). Some of the clay minerals, allow us to infer the local environmental conditions in the context of this deposit. In particular, the genesis of sepiolite has been linked to highly alkaline environments (pH > 8), caused by groundwater with a high relative abundance of Mg indicative of semi-arid climatic conditions (Galán, 2006). Moreover, the mammals of the La Paz local fauna also allow reconstructing this environmental conditions since, in general, have been postulated as indicators of open savannah grasslands or arid to semi-arid environments (Tonni *et al.*, 1999; Ortiz-Jaureguizar and Cladera, 2006).

CONCLUSIONS

The bonebed in which the La Paz local fauna is included is, until now, unique in Uruguay in terms of number and concentration of skeletal remains. The taphonomic parameters of the bonebed (disarticulation, orientation, weathering, abrasion, fracturing and selection) are indicative of a deposit generated by a high energy process, such as a mudflow. The mixture of fossils with different taphonomic histories indicates that this assemblage is parautochthonous or allochthonous or a combination of both.

The sedimentological features (massive deposit, absence of sedimentary structures, poor sorting, high degree of angularity of the largest clasts and facies associations) are also consistent with a mudflow but not in a channel. This phenomenon was the final agent responsible for the deposition of the bonebed.

The La Paz local fauna is included in sediments of the Libertad Formation. This fact represents a late middle Pleistocene-early Holocene age for that formation, based on its fauna. Regarding the absolute age of the deposit, the available data (5917 and 17620 yrs BP) cannot be considered indicative of its real age, but the sedimentary features and the intense diagenesis did not allow to obtain more data to be compared with the previous ones.

The mammals of the La Paz local fauna and the presence of particular clay minerals allow us to infer, in the context of this deposit, an arid to semi-arid environment.

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APPENDIX

Appendix: Updated systematic list of the La Paz local fauna. Based on Perea *et al.* (2001); Marchesano *et al.* (2002), Alfaro and Perea (2003); Corona *et al.* (2005); Corona *et al.* (2008b); Alvarenga *et al.* (2010), and new data.

Xenarthra: Pilosa indet., Megatheriinae indet., Mylodontidae, *Lestodon* sp. Glyptodontidae indet, *Glyptodon* sp., *Doedicurus* sp., *Panochthus* sp., *Neosclerocalyptus* sp., Pampatheriidae indet., Dasypodidae, *Propaopus* sp.
 Artiodactyla: Cervidae indet., *Ozotoceros* Cf. *O. bezoarticus*, Camelidae indet., *Hemiauchenia* sp.
 Perissodactyla: Equidae indet., *Hippidion* sp., *Equus* sp.
 Notoungulata: Toxodontidae indet., *Toxodon* sp.
 Proboscidea: Gomphotheriidae indet., *Stegomastodon waringi*.
 Litopterna: Macrauchenidae indet, *Macrauchenia* cf. *M. patachonica*.
 Rodentia: Myocastoridae indet.
 Carnivora: *Arctotherium bonariense*.
 Aves: *Ciconia* sp., Phorusracidae.

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